
Water Movements in lakes during Summer Stratification; Evidence from the Distribution of Temperature in Windermere: Appendix: Oscillations in a Three-Layered Stratified Basin

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APPENDIX. OSCILLATIONS IN A THREE-LAYERED STRATIFIED BASIN*

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Consider a rectangular basin of length l and depth h which, in equilibrium, contains three horizontal layers of fluid of different densities ρ_1, ρ_2 and ρ_3 and of depths h_1, h_2 and h_3 . (The suffixes 1, 2 and 3 refer to the upper, middle and lower layers respectively.) Let rectangular co-ordinates (x, y, z) be taken with the origin in the upper surface at one end of the basin, the x -axis horizontal and the z -axis vertically downwards. The motion will be assumed to be two-dimensional and independent of the y -co-ordinate.

Since the wave-length of the motion to be considered will be large compared with the depth, the vertical acceleration may be neglected, so that the pressure equals the hydrostatic pressure. Thus, if ζ_1, ζ_2 and ζ_3 denote the downward displacement of the upper surfaces of the three layers from their equilibrium positions, the pressures p_1, p_2 and p_3 in the three layers are given by

$$\left. \begin{aligned} p_1 &= g\rho_1(z - \zeta_1), \\ p_2 &= g\rho_1(h_1 + \zeta_2 - \zeta_1) + g\rho_2(z - h_1 - \zeta_2), \\ p_3 &= g\rho_1(h_1 + \zeta_2 - \zeta_1) + g\rho_2(h_2 + \zeta_3 - h_1 - \zeta_2) + g\rho_3(z - h_2 - \zeta_3). \end{aligned} \right\} \quad (\text{A } 1)$$

If u_1, u_2 and u_3 denote the horizontal velocities, and if squares of the velocity are neglected, we have

$$\frac{du_1}{dt} = -\frac{1}{\rho_1} \frac{dp_1}{dx}, \quad \frac{du_2}{dt} = -\frac{1}{\rho_2} \frac{dp_2}{dx}, \quad \frac{du_3}{dt} = -\frac{1}{\rho_3} \frac{dp_3}{dx}, \quad (\text{A } 2)$$

which shows that u_1, u_2 and u_3 are independent of z , i.e. that the motion in any vertical plane is uniform for each layer. Thirdly, by considering the flux of water into any element of volume we obtain the equations of continuity:

$$\frac{d}{dx} (h_1 u_1) = \frac{d}{dt} (\zeta_1 - \zeta_2), \quad \frac{d}{dx} (h_2 u_2) = \frac{d}{dt} (\zeta_2 - \zeta_3), \quad \frac{d}{dx} h_3 u_3 = \frac{d\zeta_3}{dt}. \quad (\text{A } 3)$$

On eliminating p_1, p_2, p_3, u_1, u_2 and u_3 from equations (A 1, 2 and 3) we find

$$\left. \begin{aligned} \frac{d^2}{dt^2} (\zeta_1 - \zeta_2) &= gh_1 \frac{d^2}{dx^2} [(\zeta_1 - \zeta_2) + (\zeta_2 - \zeta_3) + \zeta_3], \\ \frac{d^2}{dt^2} (\zeta_2 - \zeta_3) &= gh_1 \frac{d^2}{dx^2} \left[\frac{\rho_1}{\rho_2} (\zeta_1 - \zeta_2) + (\zeta_2 - \zeta_3) + \zeta_3 \right], \\ \frac{d^2}{dt^2} \zeta_3 &= gh_1 \frac{d^2}{dx^2} \left[\frac{\rho_1}{\rho_3} (\zeta_1 - \zeta_2) + \frac{\rho_1}{\rho_2} (\zeta_2 - \zeta_3) + \zeta_3 \right]. \end{aligned} \right\} \quad (\text{A } 4)$$

We seek a simple harmonic motion of wave-length $2\pi/k$ and period $2\pi/\sigma$; the operators d^2/dt^2 and d^2/dx^2 are therefore to be replaced by $-\sigma^2$ and $-k^2$ respectively, and we have

$$\left. \begin{aligned} (h_1 + H) (\zeta_1 - \zeta_2) + h_1 (\zeta_2 - \zeta_3) + h_1 \zeta_3 &= 0, \\ h_2 \frac{\rho_1}{\rho_2} (\zeta_1 - \zeta_2) + (h_2 + H) (\zeta_2 - \zeta_3) + h_2 \zeta_3 &= 0, \\ h_3 \frac{\rho_1}{\rho_3} (\zeta_1 - \zeta_2) + h_3 \frac{\rho_2}{\rho_3} (\zeta_2 - \zeta_3) + (h_3 + H) \zeta_3 &= 0, \end{aligned} \right\} \quad (\text{A } 5)$$

where

$$H = -\sigma^2/gk^2. \quad (\text{A } 6)$$

* Some of the following results have been obtained by Makkaweev (1936) and other writers; they are collected here for convenience of reference.

The elimination of $(\zeta_1 - \zeta_2)$, $(\zeta_2 - \zeta_3)$ and ζ_3 from (A 5) gives

$$\begin{vmatrix} 1 + \frac{H}{h_1} & 1 & 1 \\ \frac{\rho_1}{\rho_2} & 1 + \frac{H}{h_2} & 1 \\ \frac{\rho_1}{\rho_3} & \frac{\rho_2}{\rho_3} & 1 + \frac{H}{h_3} \end{vmatrix} = 0, \quad (\text{A } 7)$$

which on expansion becomes

$$\begin{aligned} & H^3 + H^2(h_1 + h_2 + h_3) \\ & - H \left[h_2 h_3 \left(\frac{\rho_2}{\rho_3} - 1 \right) + h_1 h_3 \left(\frac{\rho_1}{\rho_3} - 1 \right) + h_1 h_2 \left(\frac{\rho_1}{\rho_2} - 1 \right) \right] \\ & + h_1 h_2 h_3 \left(\frac{\rho_1}{\rho_2} - 1 \right) \left(\frac{\rho_2}{\rho_3} - 1 \right) = 0. \end{aligned} \quad (\text{A } 8)$$

Equation (A 8) determines the relation between the period and the wave-length. Since the left-hand side is a cubic in H , there will be, for a given wave-length, three possible modes of oscillation which we may denote by the indices (1), (2) and (3). For each value of $H^{(i)}$ the ratios of the corresponding displacements $\zeta_1^{(i)}$, $\zeta_2^{(i)}$ and $\zeta_3^{(i)}$ may be found from equations (A 5). Thus we have

$$\zeta_1^{(i)} : \zeta_2^{(i)} : \zeta_3^{(i)} = H^{(i)} : h_1 + H^{(i)} : \frac{H^{(i)}(h_1 + H^{(i)}) h_3 \rho_2}{h_2 h_3 (\rho_3 - \rho_2) + H^{(i)}(h_2 \rho_3 + h_3 \rho_2)}. \quad (\text{A } 9)$$

The density differences recurring in practice are of the order of one part per thousand. Thus one solution of (8) will be given very nearly by

$$H^{(1)} = -(h_1 + h_2 + h_3),$$

that is,
$$\frac{\sigma^{(1)^2}}{k^2} = gh. \quad (\text{A } 10)$$

This is the ordinary 'surface' seiche. From (A 9) we have then

$$\zeta_1^{(1)} : \zeta_2^{(1)} : \zeta_3^{(1)} = h_1 + h_2 + h_3 : h_2 + h_3 : h_3, \quad (\text{A } 11)$$

showing that the displacements are all in the same sense and are proportional to the height above the bottom. The two 'internal' seiches, in which we are chiefly interested here, are given by the remaining roots $H^{(2)}$ and $H^{(3)}$ of equation (A 5). Since these roots are of the order of $(\rho_2/\rho_3 - 1)$ they satisfy the equation

$$\begin{aligned} & H^2(h_1 + h_2 + h_3) \\ & - H \left[h_2 h_3 \left(\frac{\rho_2}{\rho_3} - 1 \right) + h_1 h_3 \left(\frac{\rho_1}{\rho_3} - 1 \right) + h_1 h_2 \left(\frac{\rho_1}{\rho_2} - 1 \right) \right] \\ & + h_1 h_2 h_3 \left(\frac{\rho_1}{\rho_2} - 1 \right) \left(\frac{\rho_2}{\rho_3} - 1 \right) = 0 \end{aligned} \quad (\text{A } 12)$$

approximately (the term in H^3 in equation (A 5) being now negligible). The ratios of the corresponding displacements are found from (9):

$$\zeta_1^{(i)} : \zeta_2^{(i)} : \zeta_3^{(i)} = H^{(i)} : h_1 : \frac{H^{(i)} h_1 h_3}{h_2 h_3 (\rho_3/\rho_2 - 1) + H^{(i)}(h_2 + h_3)} \quad (\text{A } 13)$$

($i = 2, 3$) to the same order of approximation. Thus in the two 'internal' seiches the displacement of the free surface is very small compared with that of the two interfaces. It will be convenient to denote the ratio of the other two displacements by $\beta^{(i)}$:

$$\beta^{(i)} = \zeta_2^{(i)} / \zeta_3^{(i)} = (h_2 / H^{(i)}) (\rho_3 / \rho_2 - 1) + (h_2 / h_3 + 1). \quad (\text{A } 14)$$

Since there is no flow across the vertical walls $x = 0, l$, l must be a multiple of half a wavelength; hence

$$\left. \begin{aligned} k &= n\pi/l = nk_0, \\ \sigma^{(i)} &= k \sqrt{\{-gH^{(i)}\}} = n\sigma_0^{(i)}, \end{aligned} \right\} \quad (\text{A } 15)$$

where

$$k_0 = \pi/l, \quad \sigma_0^{(i)} = (\pi/l) \sqrt{\{-gH^{(i)}\}}. \quad (\text{A } 16)$$

In general the motion will consist of the sum of an infinite number of modes whose wavelength is given by (A 15). If the initial displacement of the free surface is small compared with that of the two interfaces, the modes of 'surface' type (corresponding to $H = H^{(1)}$) will not be excited. Hence we shall have

$$\left. \begin{aligned} \zeta_1 &= 0, \\ \zeta_2 &= \sum_{n=1}^{\infty} \cos nk_0 x (\beta^{(2)} A_n^{(2)} \cos n\sigma_0^{(2)} t + \beta^{(2)} B_n^{(2)} \sin n\sigma_0^{(2)} t) \\ &\quad + \sum_{n=1}^{\infty} \cos nk_0 x (\beta^{(3)} A_n^{(3)} \cos n\sigma_0^{(3)} t + \beta^{(3)} B_n^{(3)} \sin n\sigma_0^{(3)} t), \\ \zeta_3 &= \sum_{n=1}^{\infty} \cos nk_0 x (A_n^{(2)} \cos n\sigma_0^{(2)} t + B_n^{(2)} \sin n\sigma_0^{(2)} t) \\ &\quad + \sum_{n=1}^{\infty} \cos nk_0 x (A_n^{(3)} \cos n\sigma_0^{(3)} t + B_n^{(3)} \sin n\sigma_0^{(3)} t), \end{aligned} \right\} \quad (\text{A } 17)$$

where $A_n^{(2)}$, $A_n^{(3)}$, $B_n^{(2)}$ and $B_n^{(3)}$ are constants to be determined from the initial conditions. Suppose that when $t = 0$ we have

$$\left. \begin{aligned} \zeta_2 &= f_2(x), \quad \zeta_3 = f_3(x), \\ \frac{d\zeta_2}{dt} &= g_2(x), \quad \frac{d\zeta_3}{dt} = g_3(x). \end{aligned} \right\} \quad (\text{A } 18)$$

Then on writing $t = 0$ in (A 17), we have

$$\left. \begin{aligned} \Sigma(\beta^{(2)} A_n^{(2)} + \beta^{(3)} A_n^{(3)}) \cos nk_0 x &= f_2(x), \\ \Sigma(A_n^{(2)} + A_n^{(3)}) \cos nk_0 x &= f_3(x), \end{aligned} \right\} \quad (\text{A } 19)$$

and

$$\left. \begin{aligned} \Sigma(n\sigma_0^{(2)} \beta^{(2)} B_n^{(2)} + n\sigma_0^{(3)} \beta^{(3)} B_n^{(3)}) \cos nk_0 x &= g_2(x), \\ \Sigma(n\sigma_0^{(2)} B_n^{(2)} + n\sigma_0^{(3)} B_n^{(3)}) \cos nk_0 x &= g_3(x). \end{aligned} \right\} \quad (\text{A } 20)$$

By Fourier's theorem, equation (A 19) can be satisfied provided

$$\left. \begin{aligned} \beta^{(2)} A_n^{(2)} + \beta^{(3)} A_n^{(3)} &= \frac{2}{l} \int_0^l f_2(x) \cos nk_0 x dx, \\ A_n^{(2)} + A_n^{(3)} &= \frac{2}{l} \int_0^l f_3(x) \cos nk_0 x dx, \end{aligned} \right\} \quad (\text{A } 21)$$

and so

$$\left. \begin{aligned} A_n^{(2)} &= \frac{2}{l} \int_0^l \frac{f_2(x) - \beta^{(3)} f_3(x)}{\beta^{(2)} - \beta^{(3)}} \cos nk_0 x dx, \\ A_n^{(3)} &= \frac{2}{l} \int_0^l \frac{f_2(x) - \beta^{(2)} f_3(x)}{\beta^{(3)} - \beta^{(2)}} \cos nk_0 x dx. \end{aligned} \right\} \quad (\text{A } 22)$$

Similarly,

$$\left. \begin{aligned} n\sigma^{(2)} B_n^{(2)} &= \frac{2}{l} \int_0^l \frac{g_2(x) - \beta^{(3)} g_3(x)}{\beta^{(2)} - \beta^{(3)}} \cos nk_0 x dx, \\ n\sigma^{(3)} B_n^{(3)} &= \frac{2}{l} \int_0^l \frac{g_2(x) - \beta^{(2)} g_3(x)}{\beta^{(3)} - \beta^{(2)}} \cos nk_0 x dx. \end{aligned} \right\} \quad (\text{A } 23)$$

When the motion starts from rest, it is clear that $B_n^{(2)}$ and $B_n^{(3)}$ are zero.

The coefficients $A_n^{(2)}$, $A_n^{(3)}$, $B_n^{(2)}$ and $B_n^{(3)}$ being found from (A 22) and (A 23), the complete solution is now given by (A 17).

The horizontal velocities (u_1 , u_2 and u_3 for the top, middle and bottom layers respectively) are found from the equations of continuity (equation (A 3)). These give

$$\left. \begin{aligned} h_1 u_1 &= \sum_{n=1}^{\infty} \frac{\sigma_0^{(2)}}{k_0} \sin nk_0 x (\beta^{(2)} A_n^{(2)} \sin n\sigma_0^{(2)} t - \beta^{(2)} B_n^{(2)} \cos n\sigma_0^{(2)} t) \\ &+ \sum_{n=1}^{\infty} \frac{\sigma_0^{(3)}}{k_0} \sin nk_0 x (\beta^{(3)} A_n^{(3)} \sin n\sigma_0^{(3)} t - \beta^{(3)} B_n^{(3)} \cos n\sigma_0^{(3)} t), \end{aligned} \right\} \quad (\text{A } 24)$$

$$\left. \begin{aligned} -h_3 u_3 &= \sum_{n=1}^{\infty} \frac{\sigma_0^{(2)}}{k_0} \sin nk_0 x (A_n^{(2)} \sin n\sigma_0^{(2)} t - B_n^{(2)} \cos n\sigma_0^{(2)} t) \\ &+ \sum_{n=1}^{\infty} \frac{\sigma_0^{(3)}}{k_0} \sin nk_0 x (A_n^{(3)} \sin n\sigma_0^{(3)} t - B_n^{(3)} \cos n\sigma_0^{(3)} t), \end{aligned} \right\} \quad (\text{A } 25)$$

and

$$h_2 u_2 = -(h_1 u_1 + h_3 u_3). \quad (\text{A } 26)$$

In the plane $x = \frac{1}{2}l$, and when the motion initially starts from rest, we have

$$\left. \begin{aligned} h_1 u_1 &= \sum_{n=1}^{\infty} \frac{\sigma_0^{(2)}}{k_0} \sin \frac{1}{2}n\pi \cdot \beta^{(2)} A_n^{(2)} \sin n\sigma_0^{(2)} t \\ &+ \sum_{n=1}^{\infty} \frac{\sigma_0^{(3)}}{k_0} \sin \frac{1}{2}n\pi \cdot \beta^{(3)} A_n^{(3)} \sin n\sigma_0^{(3)} t, \end{aligned} \right\} \quad (\text{A } 27)$$

and

$$\left. \begin{aligned} -h_3 u_3 &= \sum_{n=1}^{\infty} \frac{\sigma_0^{(2)}}{k_0} \sin \frac{1}{2}n\pi \cdot A_n^{(2)} \sin n\sigma_0^{(2)} t \\ &+ \sum_{n=1}^{\infty} \frac{\sigma_0^{(3)}}{k_0} \sin \frac{1}{2}n\pi \cdot A_n^{(3)} \sin n\sigma_0^{(3)} t. \end{aligned} \right\} \quad (\text{A } 28)$$

u_2 may be found from equation (A 26) as in the general case.

REFERENCES

- Aichi, K. 1918 On the theory of internal seiches. *Proc. Tokyo phys.-math. Soc.* **9**, 464–478.
 Aichi, K. 1918 Calculation of the period of the internal seiches for various lakes. *Proc. Tokyo phys.-math. Soc.* **9**, 478–485.
 Becquerel & Breschet 1836 Procédé électro-chimique pour déterminer la température de la terre et des lacs à diverses profondeurs. *C.R. Acad. Sci., Paris*, **3**, 778–781.
 Bergsten, F. 1926 The seiches of Lake Vetter. *Geogr. Ann., Stockh.*, pp. 1–73.

- Birge, E. A. 1910 On the evidence for temperature seiches. *Trans. Wis. Acad. Sci. Arts Lett.* **16**, 1005–1016.
- Birge, E. A. 1916 The work of the wind in warming the lake. *Trans. Wis. Acad. Sci. Arts Lett.* **18**, 341–391.
- Charnock, H. 1951 Meteorology: Energy transfer between the atmosphere and the ocean. *Sci. Progr. Lond.* **39**, 80–95.
- Defant, A. 1918 Neue Methode zur Ermittlung der Eigenschwingungen (Seiches) von abgeschlossenen Wassermassen (Seen, Buchten usw.). *Ann. Hydrogr., Berl.*, **46**, 78.
- Demoll, R. 1921 Temperaturwellen (Seiches) und Planktonwellen. *Arch. Hydrobiol. Plankt.* **13**, 313–320.
- Exner, F. M. 1908a Über eigentümliche Temperaturschwankungen von eintägiger Periode im Wolfgangsee. *S.B. Akad. Wiss. Wien, IIa*, **117**, 9–26.
- Exner, F. M. 1908b Ergebnisse einiger Temperaturregistrierungen im Wolfgangsee. *S.B. Akad. Wiss. Wien, IIa*, **117**, 1295–1315.
- Exner, F. M. 1910 Zur Frage der Temperatureiseiches. I. *Petermanns Mitt.* **56**, 139.
- Exner, F. M. 1928 Über Temperatureiseiches im Lunzer See. *Ann. Hydrogr., Berl.*, pp. 14, 142.
- Flück, H. 1927 Beiträge zur Kenntnis des Phytoplanktons des Brienersees. *Z. Hydrol., etc., Aarau*, **4**, 1–70.
- Forel, F. A. 1895 *Le Léman. Monographie limnologique*, **2**, 651. Lausanne: Rouge.
- Goldstein, S. 1931 On the stability of superposed streams of fluids of different densities. *Proc. Roy. Soc. A*, **132**, 524–548.
- Halbfass, W. 1909 Zur Frage der Temperatureiseiches. *Petermanns Mitt.* **55**.
- Halbfass, W. 1910a Zur Frage der Temperatureiseiches. II. *Petermanns Mitt.* **56**, 139.
- Halbfass, W. 1910b Gibt es im Madüsee Temperatureiseiches. *Int. Rev. Hydrobiol.* **3**, 1–40.
- Hellström, B. 1941 Wind effect on lakes and rivers. *IngenVetenskAkad. Handl.* **158**, 191 pp.
- Hutchinson, G. E. 1941 Limnological studies in Connecticut. IV. The mechanisms of intermediary metabolism in stratified lakes. *Ecol. Monogr.* **11**, 21–60.
- Johnsson, O. H. 1946 Termisk-hydrologiska Studier i Sjön Klämningen. *Geogr. Ann., Stockh.*, **1946**, 1–154.
- Johnsson, O. H. 1948 Wind currents in Lake Klämningen. *Union geod. geophys. int. VIIIème Ass. gen., Oslo*, **1**, 367–372.
- Makkaweev, W. M. 1936 Bemerkungen über die Theorie der langen internen Wellen. *V. Hydrologische Konferenz der Baltischen Staaten, Helsinki*, June 1936, Bericht 15C, pp. 13–19.
- Mercanton, P. L. 1932 Etude de la circulation des eaux du Lac Léman. *Mem. Soc. vaud. Sci. nat.* **4**, 225–271.
- Möller, L. 1928 Hydrographische Arbeiten am Sakrower See bei Potsdam. *Z. Ges. Erdkunde*, Sonderband zur Jahrhundertfeier, 535–51.
- Montgomery, R. B. *et al.* 1946 (Symposium) Convection patterns in atmosphere and ocean. *Ann. N.Y. Acad. Sci.* **48**, 705–844.
- Mortimer, C. H. 1941 The exchange of dissolved substances between mud and water in lakes. I. The distribution of some physical variables and concentrations of dissolved substances in Esthwaite Water, April 1939–February 1940. *J. Ecol.* **29**, 280–312.
- Mortimer, C. H. 1942 The exchange of dissolved substances between mud and water in lakes. IV. General discussion. *J. Ecol.* **30**, 168–201.
- Mortimer, C. H. 1951 The use of models in the study of water movements in stratified lakes. *Verh. internat. Vereinig. Limnologie (Proc. Int. Limnol. Ass., Stuttgart)*, **11**, 254–60.
- Munk, W. H. 1941 Internal waves in the Gulf of California. *J. Mar. Res., Yale*, **4**, 81–91.
- Munk, W. H. & Anderson, E. R. 1948 Notes on a theory of the thermocline. *J. Mar. Res., Yale*, **7**, 276–295.
- Murray, J. 1888 On the effect of winds on the distribution of temperature in the sea and fresh water lochs of the west of Scotland. *Scot. Geogr. Mag.* **4**, 345–65.

- Proudman, J. 1914 Free and forced longitudinal tidal motion in a lake. *Proc. Lond. Math. Soc.* (2), **14**, 240–50.
- Proudman, J. & Doodson, A. T. 1924 Time-relations in meteorological effects on the sea. *Proc. Lond. Math. Soc.* **24**, 140–9.
- Revelle, R. R. 1939 Sediments of the Gulf of California. *Bull. Geol. Soc. Amer.* **50**, 1929.
- Richter, E. 1897 Seenstudien. *Pencks Geogr. Abhandl. Wien*, **6**, 121–191.
- Ruttner, F. 1940 *Grundriss der Limnologie*. 167 pp. Berlin: de Gruyter.
- Saunders, J. T. & Ullyott, P. 1937 Thermo-electric apparatus for limnological research. *Int. Rev. Hydrobiol.* **34**, 572–577.
- Schmidt, W. 1908 Stehende Schwingungen in der Grenzschicht zweier Flüssigkeiten. *S.B. Akad. Wiss. Wien*, IIa, **117**, 91–102.
- Spilhaus, A. F. 1938 A bathythermograph. *J. Mar. Res., Yale*, **1**, 95–100.
- Städler, M. 1934 Untersuchungen über die hygienisch bedeutungsvollen Strömungsvorgänge im Zürichsee besonders im untern Seebecken. Solothurn. 56 pp. + 18 pl. Dissertation No. 795, Eidgen. Tech. Hochschule, Zürich.
- Sverdrup, H. U., Johnson, M. W. & Fleming, R. H. 1942 *The Oceans*. 1060 pp. New York: Prentice-Hall.
- Taylor, G. I. 1931a Effect of variation in density on the stability of superposed streams of fluid. *Proc. Roy. Soc. A*, **132**, 499–523.
- Taylor, G. I. 1931b Internal waves and turbulence in a fluid of variable density. *Rapp. Cons. Explor. Mer*, **76**, 35–43.
- Thomas, E. A. 1949 Sprungschichtneigung im Zürichsee durch Sturm. *Z. Hydrol. etc., Bale* (formerly Aarau), **11**, 527–545.
- Thomas, E. A. 1950 Auffällige biologische Folgen von Sprungschichtneigungen im Zürichsee. *Z. Hydrol., etc., Bale* (formerly Aarau), **12**, 1–24.
- Thoulet, J. 1894 Contribution à l'Etude des lacs des Vosges. *Geographie (Bull. Soc. Geogr. Paris)*, **15**.
- Ule, W. 1901 Der Würmsee (Starnbergersee) in Oberbayern. *Wiss. Veröff. Ver. Erdk., Lpz.*, **5**, 211 pp. + 8 pl.
- Wasmund, E. 1927–8 Die Strömungen im Bodensee, verglichen mit bisher in Binnenseen bekannten Strömungen. *Int. Rev. Hydrobiol.* **18**, 84–114, 231–260 (1927), **19**, 21–155 (1928).
- Watson, E. R. 1904 Movements of the waters of Loch Ness as indicated by temperature observations. *Geogr. J.* **24**, 430–437.
- Wedderburn, E. M. 1907 An experimental investigation of the temperature changes occurring in fresh-water lochs. *Proc. Roy. Soc. Edinb.* **28**, 2–20.
- Wedderburn, E. M. 1909 Temperature observations in Loch Garry (Invernesshire). With notes on currents and seiches. *Proc. Roy. Soc. Edinb.* **29**, 98–135.
- Wedderburn, E. M. 1910 Current observations in Loch Garry. *Proc. Roy. Soc. Edinb.* **30**, 312–323.
- Wedderburn, E. M. 1911 The Temperature seiche. I. Temperature observations in Madüsee Pomerania. II. Hydrodynamical theory of temperature oscillations in lakes. III. Calculation of the period of the temperature seiche in the Madüsee. *Trans. Roy. Soc. Edinb.* **47**, 619–636.
- Wedderburn, E. M. 1912 Temperature observations in Loch Earn, with a further contribution to the hydrodynamical theory of the temperature seiche. *Trans. Roy. Soc. Edinb.* **48**, 629–695.
- Wedderburn, E. M. & Watson, W. 1909 Observations with a current meter in Loch Ness. *Proc. Roy. Soc. Edinb.* **29**, 619–647.
- Wedderburn, E. M. & Williams, A. M. 1911 The temperature seiche. IV. Experimental verification of the hydrodynamical theory of temperature seiches. *Trans. Roy. Soc. Edinb.* **47**, 636–642.
- Wedderburn, E. M. & Young, A. W. 1915 Temperature observations in Loch Earn. Part II. *Trans. Roy. Soc. Edinb.* **50**, 741–767.
- Welch, P. S. 1935 *Limnology*. 471 pp. New York: McGraw-Hill.